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APRIL 7.

MR. THEODORE D. RAND in the Chair.

Twenty-five persons present.

The Serpentine of Eastern Pennsylvania.—THEODORE D. RAND called attention to the specimens of serpentine presented this evening. They had been collected from numerous localities in southeastern Pennsylvania. He regarded them, as stated in a paper read before the Academy, as belonging to at least two groups: one bordering the ancient gneiss; the other, which he believed to be much more recent, occurring in the mica schists and gneisses.

The former are altered igneous rocks, either pyroxenic or chrysotitic, the chief material being enstatite, found often but slightly altered; the latter of more doubtful and perhaps varied origin, determination of which will require much more study of thin sections under the microscope.

The bright yellow serpentine from Easttown Township, Chester Co., is probably altered chrysolite chiefly, while that from Fritz Island, near Reading, is an altered dolomite. That from Brinton's Quarry, near West Chester, contains bronzite, not entirely changed.

The Radnor serpentine is chiefly altered enstatite, but specimens presented show, also, a change from asbestos into serpentine.

No rock is better suited than serpentine to show that minerals have a life history, that they are not the unchangeable substances commonly supposed, for serpentine seems to be a stage in the life of many minerals of which magnesia is a large component, while serpentine, in its turn, decomposes into soil, or occasionally, indeed in this region frequently, into quartz.

Perido-Steatite and Diabase.—DR. FLORENCE BASCOM stated that she had recently made examination of thin sections from the serpentine of the belt running northeast and southwest from Chestnut Hill through the soapstone quarry to a point northeast of Bryn Mawr, and also of the trap of the Conshohocken dyke.

The serpentine was from the quarries on the Black Rock road, between Mill Creek and the Roberts road. The belt lies wholly within the mica schists on the southeast side of the Pre-Cambrian gneiss. The serpentine proved to be derived from a peridotite and not from a dolomite or from an enstatite rock, as in other cases mentioned. The thin sections show olivine grains with the characteristic alteration to serpentine on their peripheries; much talc or steatite is present. The rock is, therefore, a perido-steatite. The dark green crystals, conspicuous in the hand specimens, often twinned, are pseudomorphs after olivine, and not after staurolite, the forms of each resembling the other closely.